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EXAMINER

JONES, HUGH M

ART UNIT	PAPER NUMBER
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2128

SHORTENED STATUTORY PERIOD OF RESPONSE	MAIL DATE	DELIVERY MODE
3 MONTHS	01/10/2007	PAPER

Please find below and/or attached an Office communication concerning this application or proceeding.

If NO period for reply is specified above, the maximum statutory period will apply and will expire 6 MONTHS from the mailing date of this communication.

Office Action Summary

Application No.

10/016,437

Applicant(s)

DUTTA ET AL.

Examiner

Hugh Jones

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-- The MAILING DATE of this communication appears on the cover sheet with the correspondence address --

Period for Reply

A SHORTENED STATUTORY PERIOD FOR REPLY IS SET TO EXPIRE 3 MONTH(S) OR THIRTY (30) DAYS, WHICHEVER IS LONGER, FROM THE MAILING DATE OF THIS COMMUNICATION.

- Extensions of time may be available under the provisions of 37 CFR 1.136(a). In no event, however, may a reply be timely filed after SIX (6) MONTHS from the mailing date of this communication.
- If NO period for reply is specified above, the maximum statutory period will apply and will expire SIX (6) MONTHS from the mailing date of this communication.
- Failure to reply within the set or extended period for reply will, by statute, cause the application to become ABANDONED (35 U.S.C. § 133). Any reply received by the Office later than three months after the mailing date of this communication, even if timely filed, may reduce any earned patent term adjustment. See 37 CFR 1.704(b).

Status

- 1) ☒ Responsive to communication(s) filed on 06 September 2006.
- 2a) ☐ This action is **FINAL**. 2b) ☒ This action is non-final.
- 3) ☐ Since this application is in condition for allowance except for formal matters, prosecution as to the merits is closed in accordance with the practice under *Ex parte Quayle*, 1935 C.D. 11, 453 O.G. 213.

Disposition of Claims

- 4) ☒ Claim(s) 1-27 is/are pending in the application.
- 4a) Of the above claim(s) _____ is/are withdrawn from consideration.
- 5) ☐ Claim(s) _____ is/are allowed.
- 6) ☒ Claim(s) 1-27 is/are rejected.
- 7) ☐ Claim(s) _____ is/are objected to.
- 8) ☐ Claim(s) _____ are subject to restriction and/or election requirement.

Application Papers

- 9) ☐ The specification is objected to by the Examiner.
- 10) ☒ The drawing(s) filed on 10 December 2001 is/are: a) ☐ accepted or b) ☒ objected to by the Examiner.
- Applicant may not request that any objection to the drawing(s) be held in abeyance. See 37 CFR 1.85(a).
- Replacement drawing sheet(s) including the correction is required if the drawing(s) is objected to. See 37 CFR 1.121(d).
- 11) ☐ The oath or declaration is objected to by the Examiner. Note the attached Office Action or form PTO-152.

Priority under 35 U.S.C. § 119

- 12) ☐ Acknowledgment is made of a claim for foreign priority under 35 U.S.C. § 119(a)-(d) or (f).
- a) ☐ All b) ☐ Some * c) ☐ None of:
1. ☐ Certified copies of the priority documents have been received.
 2. ☐ Certified copies of the priority documents have been received in Application No. _____.
 3. ☐ Copies of the certified copies of the priority documents have been received in this National Stage application from the International Bureau (PCT Rule 17.2(a)).

* See the attached detailed Office action for a list of the certified copies not received.

Attachment(s)

- 1) ☒ Notice of References Cited (PTO-892)
- 2) ☐ Notice of Draftsperson's Patent Drawing Review (PTO-948)
- 3) ☐ Information Disclosure Statement(s) (PTO/SB/08)
Paper No(s)/Mail Date _____.
- 4) ☒ Interview Summary (PTO-413)
Paper No(s)/Mail Date. _____.
- 5) ☐ Notice of Informal Patent Application
- 6) ☐ Other: _____.

DETAILED ACTION

1. Claims 1-27 of U. S. Application 10/016,437, filed 12/10/2001, are presented for examination. Unfortunately, this action is made non-final for reasons subsequently presented. The agreements thought to have been reached during the interviews are negated by Applicant's actions in their latest response.

A third interview will not be granted. A written response will be required.

2. As recited in the interview summary (emphasis added):

The Examiner and representative conducted two lengthy interviews. The Examiner worked, in both interviews, to bring the application to issue. The Examiner agreed to withdraw the 102 rejections, including the de Kok reference because it was cumulative to the Mallick (1999)/Huffman rejections (and also because Applicants would amend around the Mallick/Huffman art. The agreement was made in order to avoid the necessity for Applicants to provide an affidavit). The Examiner also agreed to withdraw the objection to the oath.

Applicants argued at some point that some of the points raised about the Huffman art were not officially of record but appeared to agree with the Examiner about the issues. The Examiner repeatedly asked Applicant's representative to work with the Examiner in a timely manner to bring the application to issue pointing out that the Examiner was going the extra mile for Applicants and that the issues could be resolved in a timely manner. The Examiner repeatedly emphasized that the critical issue, which needed to be addressed by Applicants, was the Mallick/Huffman rejections, and that the claims must be amended around

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said art. It was the Examiner's understanding that Applicants would do so.

3. Applicants have not done so. Therefore all agreements are cancelled.

Respectfully, the Examiner expended a great deal of effort to bring this application to issue and, respectfully, is dismayed at Applicant's response.

4. Furthermore, there is apparently information regarding the teachings of Huffman that should have been provided to the office (see next section).

Information Disclosure Statement

5. Applicants are again reminded of their duties under 1.56 and 1.105. One of Applicant's main arguments has been that the ratio of compressional to shear wave velocities is not disclosed in Mallick or Huffman. The Examiner has repeatedly and unsuccessfully attempted to convey that the art does disclose said feature to Applicant.

6. The following reference has been obtained: Mallick et al. (**one of the inventors**); Shallow water flow prediction using prestack waveform inversion of conventional 3D seismic data and rock modeling; 2002; obtained from http://www.westerngeco.com/media/resources/articles/shallow_water_le.pdf. (**the assignee**)

Pages 675-676 disclose (emphasis added):

Rock properties of SWF sediments. In-situ measurements of elastic and other rock properties of SWF sediments are very limited because SWF layers are associated with very low sonic velocities. Measurement of such low velocities is difficult in a cased-hole environment (tool limitations), and open-hole logging under SWF conditions is hazardous. We therefore rely on the elastic property trends of regular sands and shales to determine the rock properties of SWF sands. In recent articles, Huffman and Castagna (2001) and Zimmer et al. (2002) reviewed the rock properties of SWF sediments based on laboratory measurements. These sediments lie in the transition zone between suspended materials in fluid and rocks around critical porosity. These nearly unconsolidated sands exhibit low bulk densities and anomalously low P- and S-wave velocities (V_P and V_S). With

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increasing pressure, they lose cohesion, causing V_s to drop faster than V_p . This causes V_p/V_s to increase and therefore the Poisson's ratio to increase. V_p/V_s of the order of 10 or higher (i.e., a Poisson's ratio of 0.49 or higher) is typical of these overpressured SWF sediments. Figure 3 summarizes some rock properties of SWF formations. Note that the high V_p/V_s and Poisson's ratio is a direct result of poor grain contact in these overpressured SWF sediments.

Seismic characterization of SWF. Seismic data has long been recognized as key to SWF detection prior to drilling. Several seismic methods are available for the detection of SWF zones. Geohazard or site-survey data provide high-resolution images at shallow depths where SWF typically occurs and are good for analyzing stratigraphic sequences. However, because these geohazard data use a very short cable length, they do not provide reliable velocity information or sufficient amplitude variation with offset (AVO) information to differentiate rock properties. Huffman and Castagna (2001) suggest using multicomponent data for SWF detection. Although multicomponent seismic data provide an accurate estimate of V_P , V_S , and density, acquiring such data is expensive, especially in deepwater where hardware limitations prevent multicomponent technology to be extended to water depths in excess of 5000 ft (1500 m). For detection of SWF and quantification of its properties, we are therefore limited to conventional three-dimensional (3D) Pwave data acquired with a long cable.

7. This was never provided in an IDS or communicated to the Examiner. This reference is material and relevant to examination of the instant application,

especially in view of Applicant's argument. In fact, Applicant's one IDS,

(obtained after much difficulty) appears mostly to just list the references already submitted by the Examiner (892 form). It also contains a search report apparently received by Applicants in 2003.

8. This is the same Huffman as the inventor Huffman, applied in the 103 rejection. Allegations that Huffman does not disclose the use of the compressional and shear velocities is, respectfully, not understood. Applicants are reminded that the col. 6, lines 6-20, of Huffman was noted in the last office action and during the two interviews. Applicants, in this response, merely *allege* (page 11, response) that the feature at issue is not present:

"Similarly, Huffman is generally directed to identifying shallow water flow hazards using marine seismic data.

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However, like Mallick 1999, Huffman does not teach or disclose comparing the pressure-wave velocity to the shear-wave velocity to determine the shallow water flow risk."

9. Said section of Huffman recites (emphasis added):

(9) As noted above, the laboratory data measurements on shear velocity could not be obtained below 400 psi. However, the predicted V_p/V_s relationship is consistent with values obtained by others for unconsolidated sand-packs. FIG. 3d, after Hamilton (1976), shows a plot of compressional and shear velocities of water saturated sands at relatively low effective stress. The abscissa 190 is the effective stress in pounds per square inch (psi) while the ordinate 192 is the velocity in meters per second (m/s). Plotted are laboratory measurements of compressional 194 and shear 196 velocities for a fine sand (grain size 0.125 to 0.149 mm.) and curves 198 and 200 for the compressional and shear velocities for a coarse sand (0.59 to 0.84 mm). The scales on both the abscissa and the ordinate are logarithmic so that at 20 psi, the V_p/V_s ratio is approximately 6.0.

(10) It is of particular interest to note that the compressional velocities for the data of FIGS. 3b and 3d the sands shows relatively little dependence upon the effective stress, and at low stresses, is approximately 2000 meters per second. The V_p/V_s ratio, on the other hand, increases from a value of about 2.5 at 1000 psi to over 6.0 at 20 psi.

(11) An effective stress of 1000 psi corresponds roughly to a subsea depth of approximately 2000 feet for normally pressured sediments. This is within the range where abnormally pressured SWF sands have been encountered in deepwater drilling. What FIGS. 3a-3d show is that if such a sand is buried and the fluid pressure builds up due to differential compaction or structural geopressuring, there is a small change in the compressional wave velocity and a large change in the shear velocity of the sand. This difference in shear wave velocity will manifest itself as a time delay, or "static" shift in the seismic data that will make the abnormally-pressured SWF sand appear thicker in time on the shear wave data due to the low shear velocities.

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Additionally, a sand with a shear velocity of 700 to 800 m/s would have a relatively small difference in shear wave impedance with an overlying clay or silt sediments whereas a sand with a shear velocity of 300 m/s or less would have a much larger difference in shear wave impedance with overlying sediments. As would be well known to those versed in the art, such a difference in shear wave impedance should be detectable by suitable seismic methods. What is important for the present invention is that the abnormal pressure in a sand body will produce a small change in compressional velocity and impedance and a large change in shear velocity and impedance: the exact magnitude of the change and the mechanism that causes the change is relatively unimportant.

Paragraphs 35-37 recite:

(35) The increase in pore pressure also causes the shear modulus and frame modulus to weaken significantly, causing a large increase in attenuation for energy propagating through the abnormally pressured sand. A review of P-wave Quality Factor (Q) data indicates that P wave Q varies between 100 at normal pressures in SWF sands to values of 10 to 20 at low effective pressures. In comparison, S-wave Q varies between 100 at normal pressures in SWF sands down to zero at very low effective pressures. These changes in the Q ratio correlate closely to the V_p/V_s ratio, so that S waves are attenuated more severely than P waves. In particular, the compressional waves that are converted to shear waves on reflectors below the sand will show severe attenuation as they try to travel upward through the abnormally pressured interval. Hence, the multicomponent data will, upon display, show severe dimming and loss of higher frequencies of the mode-converted waves that have partial ray paths through any pressured sand interval. In the present invention, this dimming may be determined by measurement of the amplitude of the reflected waves relative to a location where the dimming is not present. The loss of higher frequencies is measured by a spectral analysis of the data using known methods. These zones of severe attenuation can be correlated to the presence of the abnormally-pressured sand and can be used to map out the extent of the sand much like the AVO behavior above.

(36) Thus, reflected shear wave 434b from the deeper reflector 452 is not able to pass through the anomalous

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zone, or if it does, it is severely attenuated and delayed, so that a detector at location 434b will not pick up a reflected shear wave. Those versed in the art would recognize that the absence or delay of a reflected shear wave could also be detected on data from a single source, rather than the two different source positions illustrated in FIG. 6. However, the analysis of the data would be somewhat more difficult because of the variation in the reflection coefficient of the reflected shear wave at the interface 452 with angle of incidence.

(37) When analyzed together, the combined effects of velocity anomalies, AVO anomalies, and attenuation anomalies caused by changes in effective stress in SWF sands can be used to identify, delineate and characterize the pressure regime in these hazards. A robust analysis of these SWF sands requires seismic acquisition and processing methods that preserve amplitude and phase information, and do not distort the frequency content of the seismic signals. The anomalous signatures can then be identified using a combination of P-wave travel time tomography, coupled tomography of downgoing P-wave and upgoing S-wave events, AVO analysis, post-stack inversion, pre-stack inversion, attenuation analysis, and seismic attribute analysis. These anomalies can be mapped in 2D or 3D to delineate the extent and thickness of specific SWF zones. In cases where the data permit an accurate estimate of the velocities and moduli of the sand, a specific prediction of the pore pressure can be made using the velocity-effective pressure calibration shown in FIG. 3, calibrated to local well control information where appropriate, along with an estimate of the overburden.

10. The request for information material and relevant to the examination of this application is again repeated.

11. A 1.105 requirement for information is not being made at this time.

Drawings

12. Figure 6-8 should be designated by a legend such as --Prior Art-- because only that which is old is illustrated. See MPEP § 608.02(g). Corrected drawings

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in compliance with 37 CFR 1.121(d) are required in reply to the Office action to avoid abandonment of the application. The replacement sheet(s) should be labeled "Replacement Sheet" in the page header (as per 37 CFR 1.84(c)) so as not to obstruct any portion of the drawing figures. If the changes are not accepted by the examiner, the applicant will be notified and informed of any required corrective action in the next Office action. The objection to the drawings will not be held in abeyance.

Claim Rejections - 35 USC § 102

13. The following is a quotation of the appropriate paragraphs of 35 U.S.C. 102 that form the basis for the rejections under this section made in this Office action:

A person shall be entitled to a patent unless –

(a) the invention was known or used by others in this country, or patented or described in a printed publication in this or a foreign country, before the invention thereof by the applicant for a patent.

(e) the invention was described in (1) an application for patent, published under section 122(b), by another filed in the United States before the invention by the applicant for patent or (2) a patent granted on an application for patent by another filed in the United States before the invention by the applicant for patent, except that an international application filed under the treaty defined in section 351(a) shall have the effects for purposes of this subsection of an application filed in the United States only if the international application designated the United States and was published under Article 21(2) of such treaty in the English language.

14. **Claims 1-5, 7-27 are rejected under 35 U.S.C. 102(a) as being clearly anticipated by de Kok (PCT search report, 2003; including one of the inventors.) or Huffman et al. ("H1") (The petrophysical basis for shallow-water flow prediction using multicomponent seismic data; THE LEADING EDGE SEPTEMBER 2001; pp. 1030-1036).**

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15. Claims 1-5, 7-27 are rejected under 35 U.S.C. 102(e) as being clearly anticipated by Huffman ("H2") (6,694,261).

16. de Kok ("dk") or Huffman 1 or Huffman 2 disclose prestack waveform inversion using a genetic algorithm including:

a method for determining shallow water flow risk using seismic data comprising:

processing the seismic data to enhance its stratigraphic resolution (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

selecting a control location comprising:

performing a stratigraphic analysis on the seismic data (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035); and

evaluating the seismic attributes of the seismic data (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

applying a pre-stack waveform inversion on the seismic data at a selected control location to provide an elastic model, wherein the elastic model comprises pressure-wave velocity and shear-wave velocity (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6,

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lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035); and

applying a post-stack inversion on the seismic data using the elastic model ; and determining the shallow water flow risk using the post-stack inverted elastic model to compare the pressure-wave velocity to the shear-wave velocity (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035).

wherein the pre-stack waveform inversion comprises using a genetic algorithm comprising:

generating a plurality of elastic earth models (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

generating pre-stack synthetic seismograms for the elastic earth models (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

matching the generated seismograms with the seismic data (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

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generating a fitness for the elastic earth models (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

genetically reproducing the elastic earth models using the fitness for the elastic earth (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035), and

determining convergence of the reproduced elastic earth models to select the elastic model (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

wherein the elastic model further comprises attribute of Poisson's ratio (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

wherein the control location comprises a plurality of control locations (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

wherein performing the stratigraphic analysis comprises using the model to identify a geologic feature (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11,

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line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

wherein evaluating seismic attributes comprises using AVO techniques (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col, 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035).

a method for determining shallow water flow risk using seismic data comprising:

processing the seismic data to enhance its stratigraphic resolution (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col, 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

selecting a control location comprising:

performing a stratigraphic analysis on the seismic data (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col, 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035); and

evaluating the seismic attributes of the seismic data (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col, 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

applying a pre-stack waveform inversion on the seismic data at a selected control location to provide an elastic model, wherein the elastic model comprises

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pressure-wave velocity and shear-wave velocity (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035); and

applying a post-stack inversion on the seismic data using the elastic model; and determining the shallow water flow risk using the post-stack inverted elastic model to compare the pressure-wave velocity to the shear-wave velocity (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035).

wherein the pre-stack waveform inversion comprises using a genetic algorithm comprising:

generating a plurality of elastic earth models (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

generating pre-stack synthetic seismograms for the elastic earth models (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

matching the generated seismograms with the seismic data (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7;

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H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

generating a fitness for the elastic earth models (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035);

genetically reproducing the elastic earth models using the fitness for the elastic earth models (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035), and

determining convergence of the reproduced elastic earth models to select the elastic model (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

wherein processing the seismic data comprises using an algorithm with amplitude preserving flow (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

wherein the elastic model further comprises attribute of Poisson's ratio (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

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wherein the control location comprises a plurality of control locations (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035),

wherein performing the stratigraphic analysis comprises using the model to identify a geologic feature (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035).

wherein evaluating seismic attributes comprises using AVO techniques (dk: sections entitled method, Shallow waterflow detection, and Conclusion. Note fig. 1-2, 7; H2: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49; H1: fig. 6-9; page 1033, col. 2 to first paragraph, page 1035).

Claim Rejections - 35 USC § 103

17. The following is a quotation of 35 U.S.C. 103(a) which forms the basis for all obviousness rejections set forth in this Office action:

(a) A patent may not be obtained though the invention is not identically disclosed or described as set forth in section 102 of this title, if the differences between the subject matter sought to be patented and the prior art are such that the subject matter as a whole would have been obvious at the time the invention was made to a person having ordinary skill in the art to which said subject matter pertains. Patentability shall not be negated by the manner in which the invention was made.

18. The factual inquiries set forth in *Graham v. John Deere Co.*, 383 U.S. 1, 148 USPQ 459 (1966), that are applied for establishing a background for determining obviousness under 35 U.S.C. 103(a) are summarized as follows:

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1. Determining the scope and contents of the prior art.
2. Ascertaining the differences between the prior art and the claims at issue.
3. Resolving the level of ordinary skill in the pertinent art.
4. Considering objective evidence present in the application indicating obviousness or nonobviousness.

19. Claims 1-5, 7-27 are rejected under 35 U.S.C. 103(a) as being unpatentable over Mallick (3/1999) in view of Huffman (6,694,261).

20. Mallick discloses all limitations, as subsequently discussed, but does not expressly disclose the application of the technique to Shallow Water Flow (SWF).

21. Huffman discloses a method for identification of shallow water flow hazards using seismic data (see title), using the same types of techniques.

23. It would have been obvious to one of ordinary skill in the art at the time of the invention to modify the Mallick teaching to include the Huffman teaching because Huffman disclose in the "background of the art" that there is a need to identify SWF prior to drilling a borehole.

24. Specifically, Mallick/Huffman discloses:

a method for determining shallow water flow risk using seismic data comprising ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49):

processing the seismic data to enhance its stratigraphic resolution ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49);

selecting a control location comprising:

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performing a stratigraphic analysis on the seismic data ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49); and

evaluating the seismic attributes of the seismic data ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49);

applying a pre-stack waveform inversion on the seismic data at a selected control location to provide an elastic model, wherein the elastic model comprises pressure-wave velocity and shear-wave velocity ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49); and

applying a post-stack inversion on the seismic data using the elastic model; and determining the shallow water flow risk using the post-stack inverted elastic model to compare the pressure-wave velocity to the shear-wave velocity ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49).

wherein the pre-stack waveform inversion comprises using a genetic algorithm ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49) comprising:

generating a plurality of elastic earth models ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49),

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generating pre-stack synthetic seismograms for the elastic earth models ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49);

matching the generated seismograms with the seismic data ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49);

generating a fitness for the elastic earth models ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49);

genetically reproducing the elastic earth models using the fitness for the elastic earth models ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49), and

determining convergence of the reproduced elastic earth models to select the elastic model ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49),

wherein processing the seismic data comprises using an algorithm with amplitude preserving flow ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49),

wherein the elastic model further comprises attribute of Poisson's ratio (pg. 330, col. 2; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49),

wherein the control location comprises a plurality of control locations (pg. 330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49),

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wherein performing the stratigraphic analysis comprises using the model to identify a geologic feature ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49).

wherein evaluating seismic attributes comprises using AVO techniques ("Genetic Algorithm – a Practical Implementation – pp. 326-330; H: col. 6, lines 6-20, col. 6; col. 11, line 60 to col. 12, line 49).

20. Claim 6 is rejected under 35 U.S.C. 103(a) as being unpatentable over Mallick (3/1999) in view of Huffman (6,694,261) and in further view of Tygel et al. or de Kok in view of Tygel et al..

25. Mallick or de Kok discloses all limitations, as discussed, but do not expressly disclose the use of amplitude preserving techniques.

26. Tygel et al. discloses the use of amplitude preserving techniques (page 945, top of middle column).

27. It would have been obvious to one of ordinary skill in the art at the time of the invention to modify the Mallick or de Kok teaching to incorporate the Tygel et al. teaching because Tygel et al. disclose that the use of amplitude preserving techniques reduce the deleterious effects of aliasing (page 945, top of middle column), in the same context.

Response to Arguments

28. Applicant's arguments, filed 9/6/2006 have been carefully considered, but are not persuasive.

29. Applicants arguments relating to the 103 rejections are not persuasive. Applicants are reminded that the Poisson's ratio is the well known ratio between

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pressure and shear waves. This ratio is disclosed in both references. For example, see left hand column, top of page 329 of Mallick (*one of the inventors*) and col. 6, lines 6-20, col. 6 of Huffman.

30. Applicants have recognized as much. Mallick et al. (**one of the inventors**); Shallow water flow prediction using prestack waveform inversion of conventional 3D seismic data and rock modeling; 2002; obtained from http://www.westerngeco.com/media/resources/articles/shallow_water_le.pdf. (**the assignee**)

Pages 675-676 disclose (emphasis added):

Rock properties of SWF sediments. In-situ measurements of elastic and other rock properties of SWF sediments are very limited because SWF layers are associated with very low sonic velocities. Measurement of such low velocities is difficult in a cased-hole environment (tool limitations), and open-hole logging under SWF conditions is hazardous. We therefore rely on the elastic property trends of regular sands and shales to determine the rock properties of SWF sands. In recent articles, Huffman and Castagna (2001) and Zimmer et al. (2002) reviewed the rock properties of SWF sediments based on laboratory measurements. These sediments lie in the transition zone between suspended materials in fluid and rocks around critical porosity. These nearly unconsolidated sands exhibit low bulk densities and anomalously low P- and S-wave velocities (V_P and V_S). With increasing pressure, they lose cohesion, causing V_S to drop faster than V_P . This causes V_P/V_S to increase and therefore the Poisson's ratio to increase. V_P/V_S of the order of 10 or higher (i.e., a Poisson's ratio of 0.49 or higher) is typical of these overpressured SWF sediments. Figure 3 summarizes some rock properties of SWF formations. Note that the high V_P/V_S and Poisson's ratio is a direct result of poor grain contact in these overpressured SWF sediments.

Seismic characterization of SWF. Seismic data has long been recognized as key to SWF detection prior to drilling. Several seismic methods are available for the detection of SWF zones. Geohazard or site-survey data provide high-resolution images at shallow depths where SWF typically occurs and are good for analyzing stratigraphic sequences. However, because these geohazard data use a very short cable length, they do not provide reliable velocity information or sufficient amplitude variation with offset (AVO) information to differentiate rock properties. Huffman and Castagna (2001) suggest using multicomponent data for SWF detection. Although multicomponent seismic data provide an accurate estimate of V_P , V_S , and density, acquiring such data is expensive, especially in deepwater where hardware limitations prevent multicomponent technology to be extended to water depths in excess of 5000 ft (1500 m). For detection of SWF and quantification of its properties, we are therefore limited to conventional three-dimensional (3D) Pwave data acquired with a long cable.

31. This is the same Huffman as the inventor Huffman, applied in the 103 rejection. Allegations that Huffman does not disclose the use of the compressional and shear velocities is, respectfully, not understood. Applicants

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are reminded that the col. 6, lines 6-20, of Huffman was noted in the last office action and during the two interviews. Applicants, in this response, merely *allege* (page 11, response) that the feature at issue is not present:

"Similarly, Huffman is generally directed to identifying shallow water flow hazards using marine seismic data. However, like Mallick 1999, Huffman does not teach or disclose comparing the pressure-wave velocity to the shear-wave velocity to determine the shallow water flow risk."

32. Said section of Huffman recites (emphasis added):

(9) As noted above, the laboratory data measurements on shear velocity could not be obtained below 400 psi. However, the predicted V_p/V_s relationship is consistent with values obtained by others for unconsolidated sand-packs. FIG. 3d, after Hamilton (1976), shows a plot of compressional and shear velocities of water saturated sands at relatively low effective stress. The abscissa 190 is the effective stress in pounds per square inch (psi) while the ordinate 192 is the velocity in meters per second (m/s). Plotted are laboratory measurements of compressional 194 and shear 196 velocities for a fine sand (grain size 0.125 to 0.149 mm.) and curves 198 and 200 for the compressional and shear velocities for a coarse sand (0.59 to 0.84 mm). The scales on both the abscissa and the ordinate are logarithmic so that at 20 psi, the V_p/V_s ratio is approximately 6.0.

(10) It is of particular interest to note that the compressional velocities for the data of FIGS. 3b and 3d the sands shows relatively little dependence upon the effective stress, and at low stresses, is approximately 2000 meters per second. The V_p/V_s ratio, on the other hand, increases from a value of about 2.5 at 1000 psi to over 6.0 at 20 psi.

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(11) An effective stress of 1000 psi corresponds roughly to a subsea depth of approximately 2000 feet for normally pressured sediments. This is within the range where abnormally pressured SWF sands have been encountered in deepwater drilling. What FIGS. 3a-3d show is that if such a sand is buried and the fluid pressure builds up due to differential compaction or structural geopressuring, there is a small change in the compressional wave velocity and a large change in the shear velocity of the sand. This difference in shear wave velocity will manifest itself as a time delay, or "static" shift in the seismic data that will make the abnormally-pressured SWF sand appear thicker in time on the shear wave data due to the low shear velocities. Additionally, a sand with a shear velocity of 700 to 800 m/s would have a relatively small difference in shear wave impedance with an overlying clay or silt sediments whereas a sand with a shear velocity of 300 m/s or less would have a much larger difference in shear wave impedance with overlying sediments. As would be well known to those versed in the art, such a difference in shear wave impedance should be detectable by suitable seismic methods. What is important for the present invention is that the abnormal pressure in a sand body will produce a small change in compressional velocity and impedance and a large change in shear velocity and impedance: the exact magnitude of the change and the mechanism that causes the change is relatively unimportant.

Paragraphs 35-37 recite:

(35) The increase in pore pressure also causes the shear modulus and frame modulus to weaken significantly, causing a large increase in attenuation for energy propagating through the abnormally pressured sand. A review of P-wave Quality Factor (Q) data indicates that P wave Q varies between 100 at normal pressures in SWF sands to values of 10 to 20 at low effective pressures. In comparison, S-wave Q varies between 100 at normal pressures in SWF sands down to zero at very low effective pressures. These changes in the Q ratio correlate closely to the V_p/V_s ratio, so that S waves are attenuated more severely than P waves. In particular, the compressional waves that are converted to shear waves on reflectors below the sand will show severe attenuation as they try to travel upward through the abnormally pressured interval. Hence, the multicomponent data will, upon display,

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show severe dimming and loss of higher frequencies of the mode-converted waves that have partial ray paths through any pressured sand interval. In the present invention, this dimming may be determined by measurement of the amplitude of the reflected waves relative to a location where the dimming is not present. The loss of higher frequencies is measured by a spectral analysis of the data using known methods. These zones of severe attenuation can be correlated to the presence of the abnormally-pressured sand and can be used to map out the extent of the sand much like the AVO behavior above.

(36) Thus, reflected shear wave 434b from the deeper reflector 452 is not able to pass through the anomalous zone, or if it does, it is severely attenuated and delayed, so that a detector at location 434b will not pick up a reflected shear wave. Those versed in the art would recognize that the absence or delay of a reflected shear wave could also be detected on data from a single source, rather than the two different source positions illustrated in FIG. 6. However, the analysis of the data would be somewhat more difficult because of the variation in the reflection coefficient of the reflected shear wave at the interface 452 with angle of incidence.

(37) When analyzed together, the combined effects of velocity anomalies, AVO anomalies, and attenuation anomalies caused by changes in effective stress in SWF sands can be used to identify, delineate and characterize the pressure regime in these hazards. A robust analysis of these SWF sands requires seismic acquisition and processing methods that preserve amplitude and phase information, and do not distort the frequency content of the seismic signals. The anomalous signatures can then be identified using a combination of P-wave travel time tomography, coupled tomography of downgoing P-wave and upgoing S-wave events, AVO analysis, post-stack inversion, pre-stack inversion, attenuation analysis, and seismic attribute analysis. These anomalies can be mapped in 2D or 3D to delineate the extent and thickness of specific SWF zones. In cases where the data permit an accurate estimate of the velocities and moduli of the sand, a specific prediction of the pore pressure can be made using the velocity-effective pressure calibration shown in FIG. 3, calibrated to local well control information where appropriate, along with an estimate of the overburden.

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33. Any inquiry concerning this communication or earlier

communications from the examiner should be:

directed to: Dr. Hugh Jones telephone number (571) 272-3781,

Monday-Thursday 0830 to 0700 ET,

or

the examiner's supervisor, Kamini Shah, telephone number (571) 272-2279.

Any inquiry of a general nature or relating to the status of this application should be directed to the Group receptionist, telephone number (703) 305-3900.

mailed to:

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or faxed to:

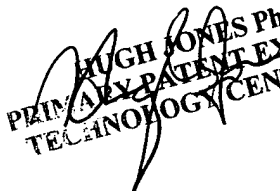
(703) 308-9051 (for formal communications intended for entry)

or (703) 308-1396 (for informal or draft communications, please label *PROPOSED* or *DRAFT*).

Dr. Hugh Jones

Primary Patent Examiner

November 24, 2006


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